

Autumn cooling paused increased CO₂ release in central Eurasia

Received: 13 May 2022

Accepted: 9 February 2023

Published online: 16 March 2023

 Check for updates

Masayuki Kondo ^{1,4}✉, Motoki Sasakawa ², Toshinobu Machida ²,
Mikhail Arshinov ³ & Tetsuya Hiyama ¹

ARISING FROM: R. Tang et al. *Nature Climate Change* <https://doi.org/10.1038/s41558-022-01304-w> (2022).

In a recent publication in *Nature Climate Change*, Tang et al.¹ reported an increase in net CO₂ release under northern autumn cooling from 2004 to 2018, indicating that both autumn warming and cooling result in net CO₂ release. Here we show that the conclusion regarding net CO₂ release under autumn cooling was impacted by the choice of the autumn period, which resulted in overlooking the appropriate cooling regions. Our analysis of individual months in autumn with empirical upscaling of eddy flux observations (FLUXCOM², the same data used by Tang et al.¹) and atmospheric CO₂ measurements from seven towers³ suggested that the increased net CO₂ release paused during the 2004–2018 autumn cooling in central Eurasia.

With ‘autumn’ defined as a period from September to November, Tang et al.¹ concluded that the widespread autumn cooling for 2004–2018 induced an increasing trend in the net CO₂ release. However, this definition of autumn is not ideal for assessing the effect of autumn cooling on CO₂ fluxes because three-month averages obscure whether the cooling occurred in months with temperatures above 0 °C. As CO₂ fluxes, particularly gross primary production (GPP), become extremely small below 0 °C, it is possible that the autumn cooling barely affected fluxes in regions where cooling occurred in months when temperatures were already below 0 °C. To evaluate this aspect, we used Climate Research Union (CRU) TS4.05 (ref. ⁴) data to classify the land north of 25° N into four regions by months when temperatures fall below 0 °C (Fig. 1a): regions with monthly mean temperature (MMT) of all three months below 0 °C (Reg1), regions with MMT of October and November below 0 °C (Reg2), regions with MMT of November below 0 °C (Reg3) and regions with MMT of all three months above 0 °C (Reg4).

As Tang et al.¹ reported, CRU TS4.05 data indicate that the cooling trend for September–October–November (SON) is widespread north of 25° N, and the cooling areas largely overlap with Reg2 and Reg3 of both North America and Eurasia (Fig. 1b). However, for September–October (SO), the spatial extent of the cooling trend reduced northward in North America; consequently, the cooling areas remained in Reg2 and barely overlapped with Reg3 (Fig. 1b). Having these results, caution should be applied while considering whether North America is an appropriate

region for evaluating the relationship between autumn cooling and CO₂ fluxes. Variations in CO₂ fluxes may be small in the cooling area overlapping with Reg2 as only sparse vegetation inhabits that area (Supplementary Fig. 1). Similarly, the cooling area in Reg3 is unlikely to cause a large variation in CO₂ fluxes because the cooling in this region occurs in November when MMT is below 0 °C (Fig. 1b). Regional mean temperatures (Fig. 1c–h) indicated that contrary to that of SON (Fig. 1d), SO temperature for Reg3 of North America increased during 2004–2018 (Fig. 1g), suggesting that the apparent cooling trend created by averaging the data for the three months concealed the warming trend.

Contrary to that in North America, the widespread cooling trend in Eurasia remained the same for SON and SO, overlapping the most in Reg3 (Fig. 1b). Regional mean SON and SO temperatures indicated that among all the regions north of 25° N, Reg2–3 of Eurasia largely reflected a transition from autumn warming to cooling around 2004 (Fig. 1c,d,f,g). As the Eurasian cooling area largely overlaps with the distribution of boreal forests, this region probably impacted CO₂ fluxes (Supplementary Fig. 1).

Assessment of the contrasting effects of the autumn cooling and warming on CO₂ fluxes indicates that monthly GPP and terrestrial ecosystem respiration (TER) from FLUXCOM tended to show positive trends for 2004–2018 in Reg2–3 of North America but a negative trend in Reg2–3 of Eurasia (Fig. 1i,j). This result was consistent with our expectation that, for North America, the warming effect on CO₂ fluxes is more profound than the cooling effect and vice versa for Eurasia; thus, trends in CO₂ fluxes should be opposite between these regions. In Reg4 of North America and Eurasia, where the autumn warming is profound (Fig. 1e,h), trends in monthly TER were positive whereas those in GPP varied monthly (Fig. 1k). In all the regions, trends in monthly TER tended to be greater than those in monthly GPP, indicating that autumn temperatures influence TER more than GPP (Fig. 1i–k). This result is opposite to the conclusion of Tang et al.¹ that the autumn cooling induced a greater decrease in GPP than in TER.

Considering GPP and TER trends, it is not the best approach to group the lands north of 25° N (that is, Reg1–4) to assess the effect of

¹Institute for Space-Earth Environmental Research, Nagoya University, Nagoya, Japan. ²National Institute for Environmental Studies, Tsukuba, Japan.

³V. E. Zuev Institute of Atmospheric Optics, Russian Academy of Sciences, Siberian Branch, Tomsk, Russia. ⁴Present address: The IDEC Institute, Hiroshima University, Higashi-Hiroshima, Japan. ✉e-mail: redmk92@gmail.com

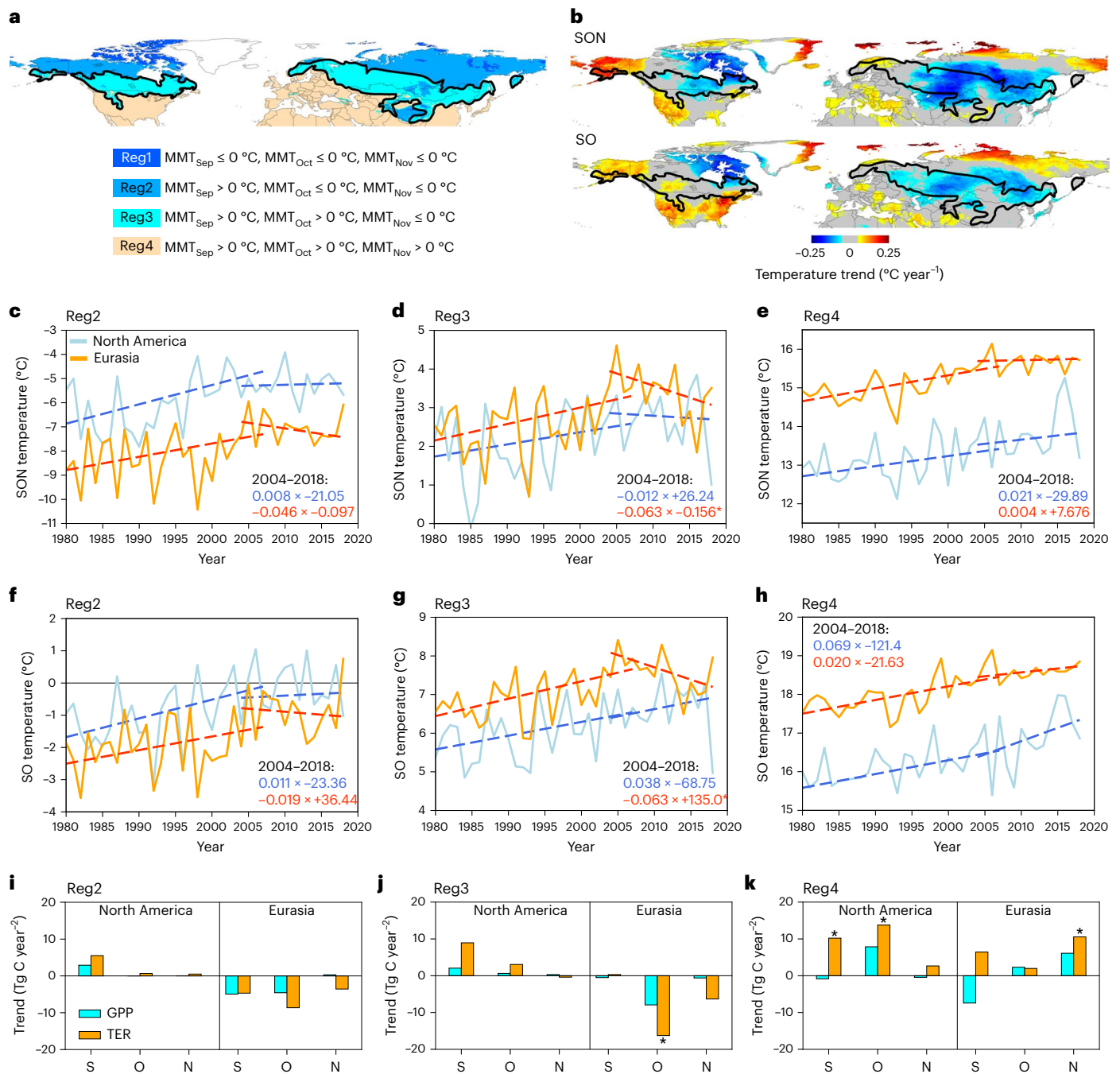


Fig. 1 | Regional patterns of 2004–2018 autumn cooling and associated GPP and TER trends for the land north of 25° N. **a**, Spatial patterns of four regions (Reg1–4) classified by MMT for 2004–2018 based on CRU TS4.05. **b**, Spatial patterns of trends in SON and SO temperatures during 2004–2018. Thick black lines in **a, b** represent a border of Reg3. **c–h**, Interannual variability in regional mean SON temperatures for Reg2 (**c**), Reg3 (**d**) and Reg4 (**e**) of North America and Eurasia, and in regional mean SO temperatures for Reg2 (**f**), Reg3 (**g**) and Reg4

(**h**) of North America and Eurasia. Linear regressions are shown for the periods 1980–2006 and 2004–2018 for North America (blue dashed lines) and Eurasia (red dashed lines), respectively, and statistics of linear regression for the period 2004–2018 are shown for North America (blue letters) and Eurasia (red letters). **i–k**, Trends of GPP and TER from FLUXCOM for Reg2 (**i**), Reg3 (**j**) and Reg4 (**k**) of North America and Eurasia for the period 2004–2018. * indicates that $P < 0.05$ for all results. S, September; O, October; N, November.

autumn cooling on net ecosystem exchange (NEE). Even with Reg4 excluded from the estimate, the autumn (defined as SON) NEE from FLUXCOM did not show a decreasing trend during 2004–2018 for all the lands north of 25° N (Fig. 2a; corresponding to Fig. 2a by Tang et al.). Focusing on the individual continent, autumn NEE for North America continuously increased in Reg1–3 since 1980 and including Reg4 enhanced the increasing trend for 2004–2018 (Fig. 2b). However, irrespective of whether Reg4 was included, autumn NEE

for Eurasia indicated a transition from an increasing to a decreasing trend around 2004 with a notable anomalous decrease between 2011 and 2016 (Fig. 2c). This is consistent with the pattern of temperatures (Fig. 1d,g).

Multi-site CO₂ measurements in central Eurasia, where significant autumn cooling occurred around the towers (Supplementary Figs. 2 and 3 and Supplementary Table 1), also showed a decrease in atmospheric CO₂ between 2011 and 2016 in response to the SO temperature, regardless

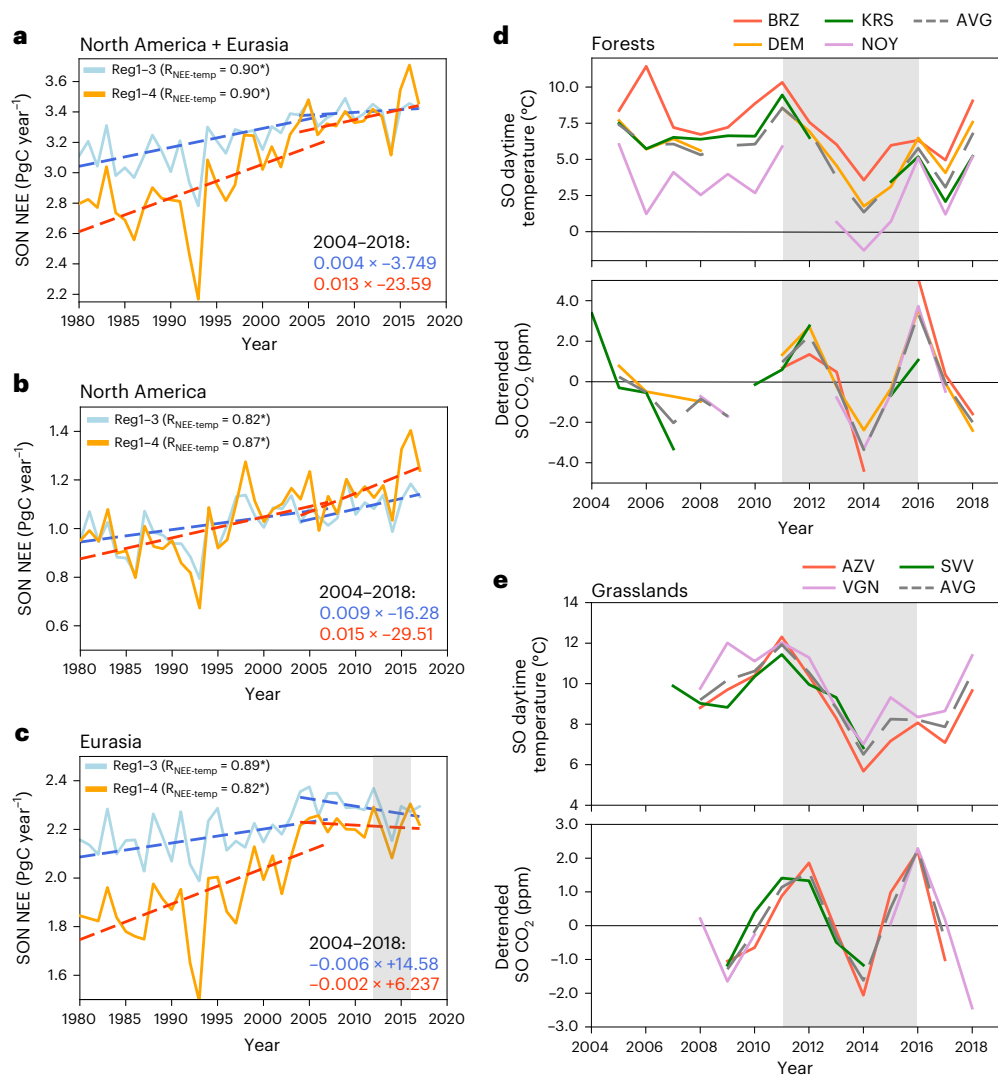


Fig. 2 | Interannual variability of autumn NEE and CO₂. **a–c**, Interannual variability in SON NEE from FLUXCOM for the lands north of 25° N (including North America and Eurasia) (**a**), North America (**b**) and Eurasia (**c**). Results are shown for the sums of Reg1–3 and Reg1–4 along with Pearson correlations between FLUXCOM NEE and CRU TS4.05 temperature (* indicates that $P < 0.05$ for all results). Linear regressions are shown for the periods 1980–2006 and 2004–2018 for Reg1–3 (blue dashed lines) and Reg1–4 (red dashed lines), respectively, and statistics of linear regression for the period 2004–2018 are

shown for Reg1–3 (blue letters) and Reg1–4 (red letters). Interannual variability in daytime temperature (12:00–18:00) and detrended CO₂ (corresponding to the same daytime) averaged for SO. **d–e**, Results are shown for towers surrounded by forests, Berezorechka (BRZ), Demyanskoe (DEM), Karasevoe (KRS) and Noyabrsk (NOY), and average (AVG) of the four site data (**d**) and grasslands, Azovo (AZV), Savvushka (SVV) and Vaganovo (VGN), and average (AVG) of the three site data (**e**) (Supplementary Table 1). Grey shading in **c–e** represents the period in which anomalous decreases in temperature and NEE were found.

of the surrounding ecosystem types: forests or grasslands (Fig. 2d,e). Tang et al.¹ used atmospheric CO₂ measurements at Point Barrow as observational support for increasing CO₂ release under autumn cooling, but this is misleading because Point Barrow is targeted to measure the background concentration in the northern ‘high’ latitudes⁵, where the autumn warming is more profound (for example, Alaska and north-eastern Siberia; Fig. 1b). Further, even if Point Barrow reflects changes in atmospheric CO₂ in northern middle latitudes, it is unlikely to detect the autumn cooling effect from its measurements as the warming effect in North America overwhelms the cooling effect in Eurasia (Fig. 2a–c).

In summary, the conclusion drawn by Tang et al.¹ could be attributed to overlooking the appropriate autumn period and cooling region. As we have demonstrated, a three-month average resulted in misidentifying the regions where the autumn cooling or warming affected CO₂ fluxes, thus leading to the biased NEE–temperature relationship. A detailed evaluation of individual months is recommended for studies to evaluate relationships between seasonal climate and CO₂ fluxes.

Online content

Any methods, additional references, Nature Portfolio reporting summaries, source data, extended data, supplementary information, acknowledgements, peer review information; details of author contributions and competing interests; and statements of data and code availability are available at <https://doi.org/10.1038/s41558-023-01625-4>.

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Reporting summary

Further information on research design is available in the Nature Portfolio Reporting Summary linked to this article.

Data availability

CRU TS4.05, FLUXCOM and atmospheric CO₂ measurements data used in this study are available respectively from the Climatic Research Unit at the University of East Anglia (https://crudata.uea.ac.uk/cru/data/hrg/cru_ts_4.05/), the data portal of the Max Planck Institute for Biogeochemistry (<https://www.bgc-jena.mpg.de/geodb/projects/Home.php>) and the Global Environmental Dataset of the Center for Global Environmental Research at the National Institute for Environmental Studies, Japan (<https://db.cger.nies.go.jp/portal/geds/atmosphericAndOceanicMonitoring?lang=eng>).

Code availability

The codes used for the analyses are available from the corresponding author.

Acknowledgements

M.K. and T.H. acknowledge support from JSPS KAKENHI (grant nos. 19H05668 and 19K12294). The measurements in central Eurasia were financially supported by the Global Environmental Research Coordination System from the Ministry of the Environment of Japan and partly by the Ministry of Science and Higher Education

of the Russian Federation (project no. 121031500342-0). M.K. acknowledges M. Takeuchi and K. Suganuma at the Graduate School of Bioagricultural Sciences, Nagoya University for supporting the processing of the Siberian CO₂ data.

Author contributions

M.K. designed the study, conducted the analysis and wrote the paper. All authors contributed to the discussion and manuscript writing.

Competing interests

The authors declare no competing interests.

Additional information

Supplementary information The online version contains supplementary material available at <https://doi.org/10.1038/s41558-023-01625-4>.

Correspondence and requests for materials should be addressed to Masayuki Kondo.

Peer review information *Nature Climate Change* thanks Minkyu Moon and the other, anonymous, reviewer(s) for their contribution to the peer review of this work.

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Research sample	The study was based on climate data (CRU-TS 4.05), empirical upscaling of eddy covariance data (FLUXCOM), and CO ₂ observations taken by seven towers in central Eurasia.
Sampling strategy	The analysis focused on the north of 25° north. The region was further classified into four regions by months when temperatures fall below 0 °C.
Data collection	All data were obtained online. All the URL information is stated in the main text.
Timing and spatial scale	The analysis and the data used were on a monthly scale, covering the period from 1980 to 2018.
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- MRI-based neuroimaging

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Antibodies used

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Validation

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